



PIERRE
AUGER
OBSERVATORY

The Importance of Atmospheric Monitoring at the Pierre Auger Observatory

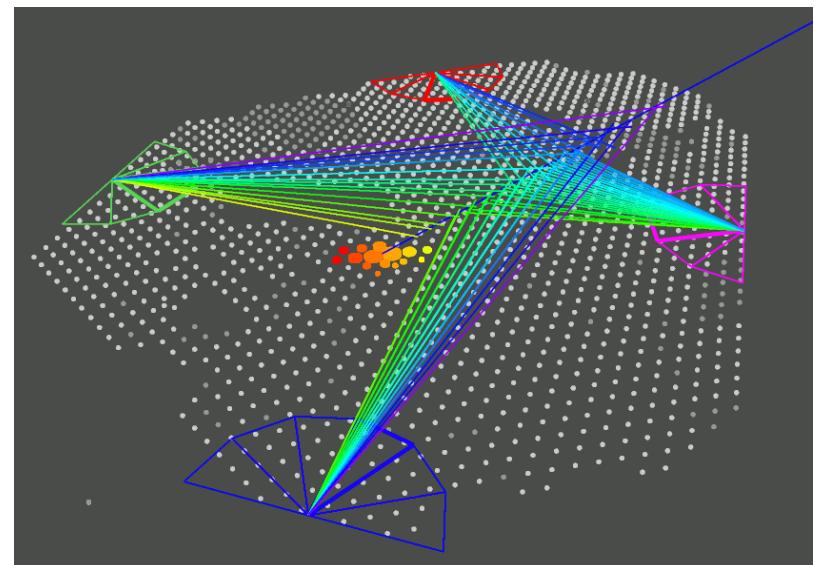
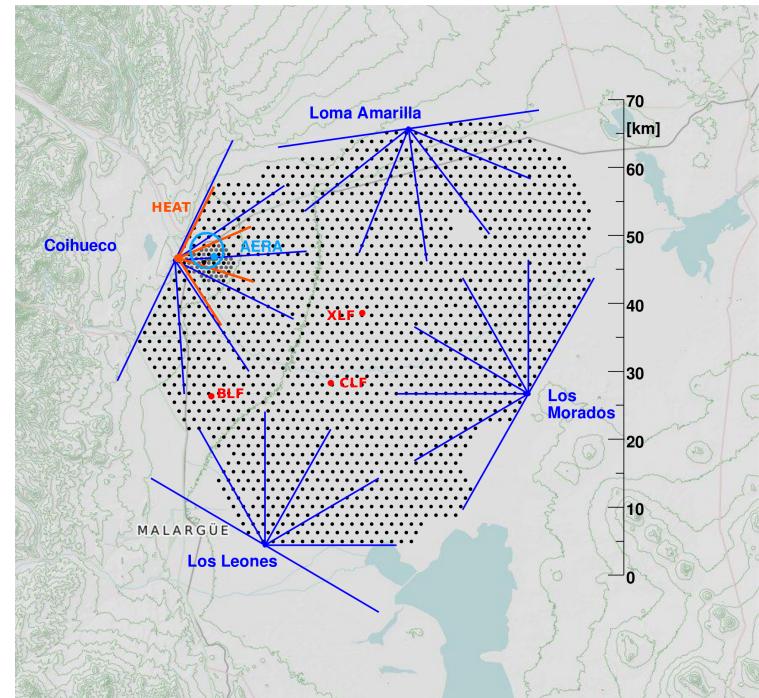
Bruce Dawson
for the Pierre Auger Collaboration

Photo: Steven Saffi, University of Adelaide

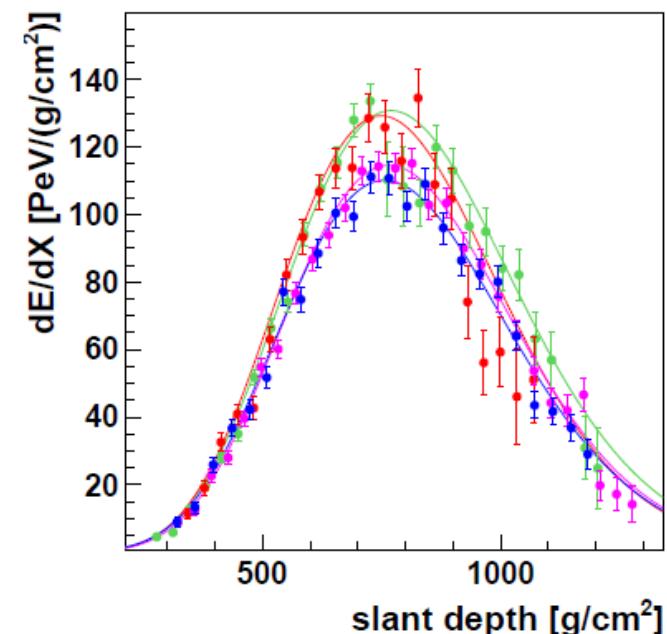
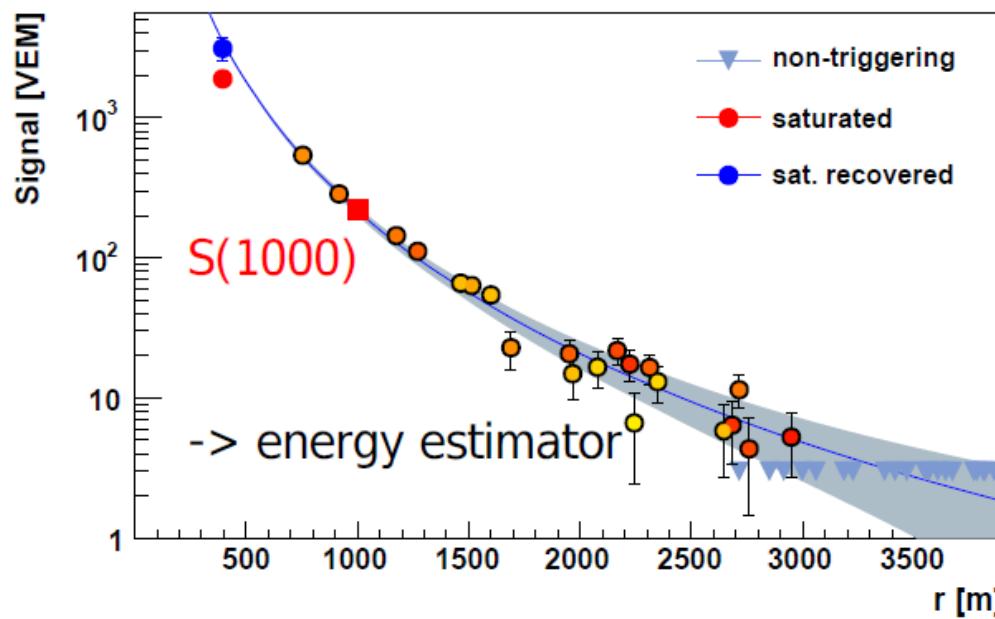
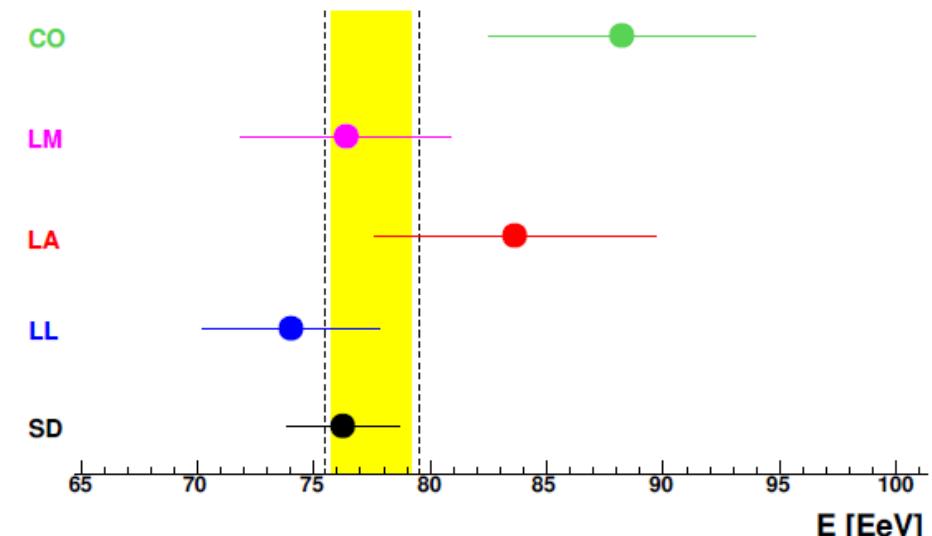
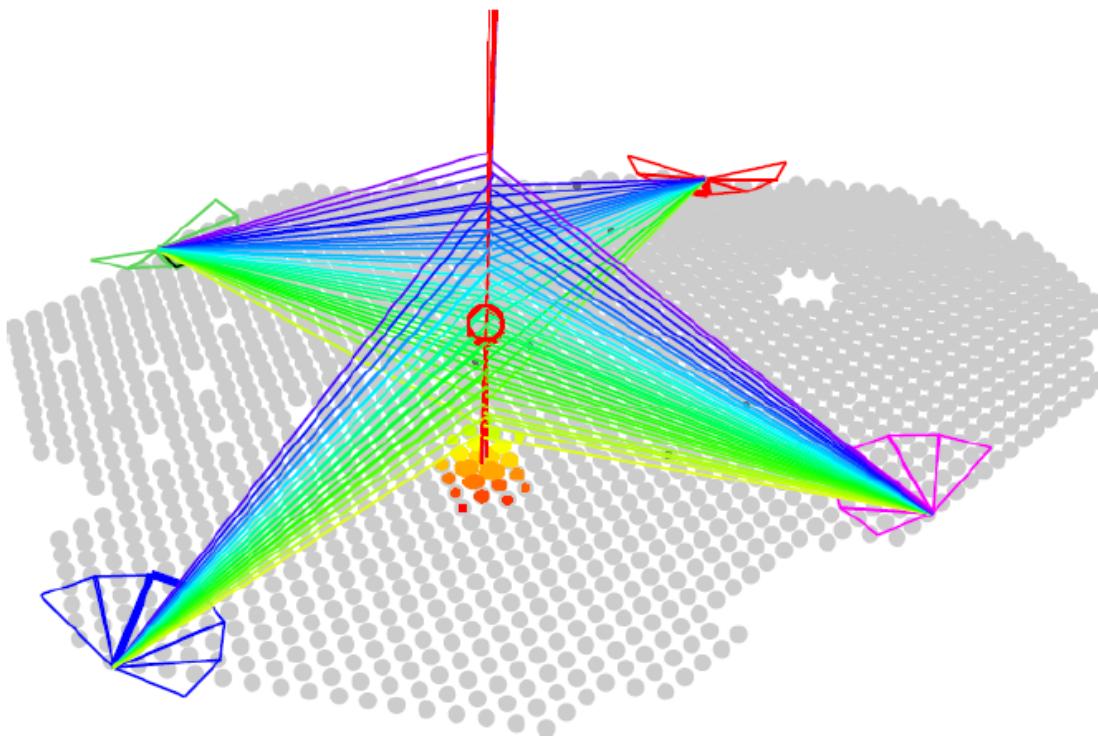


The task

- 3000 square kilometres
- surface and fluorescence detectors (SD, FD)
- FD sets energy scale of Observatory; systematic uncertainty 14%
- Most atmospheric tasks related to FD, but SD needs some atmospheric corrections

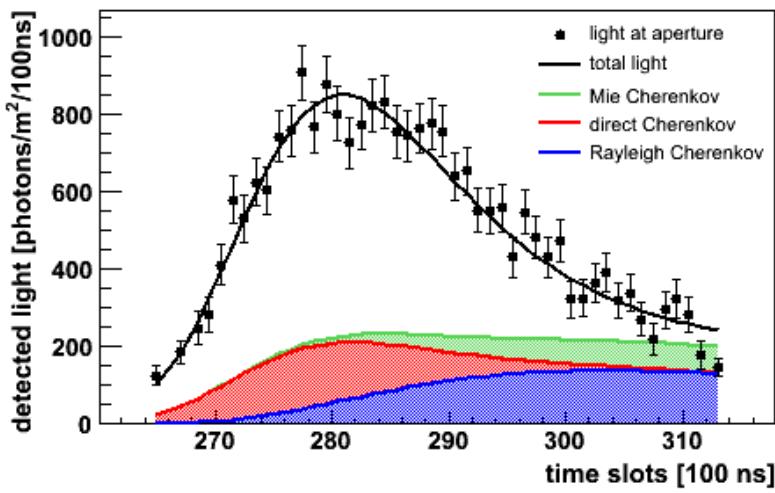
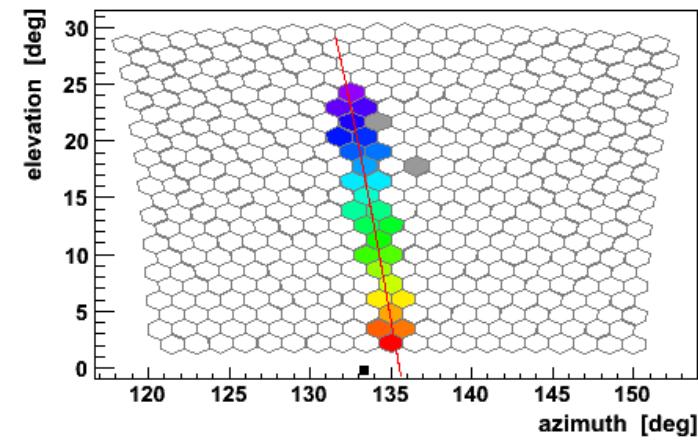
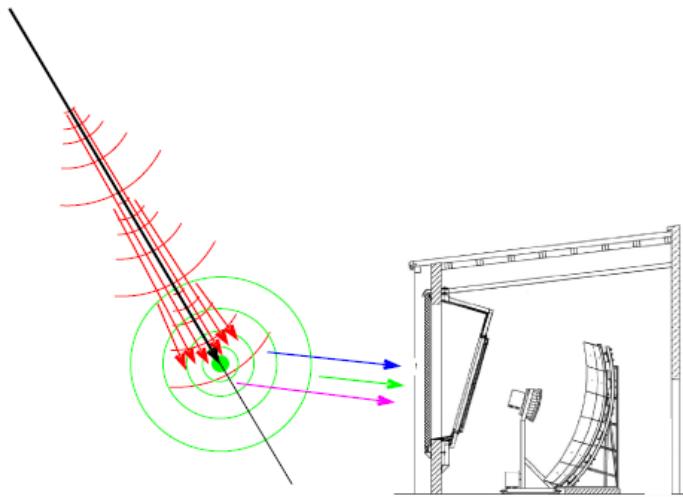


Reconstructing the air shower

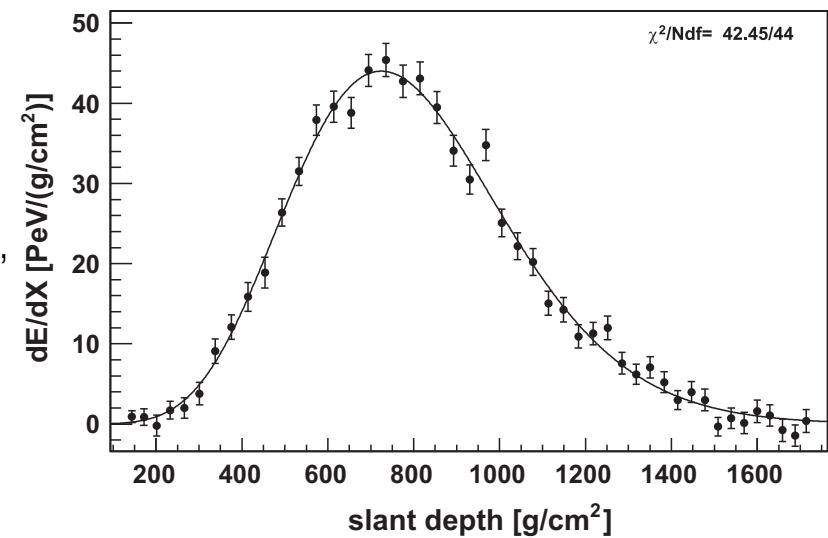


FD (hybrid) Reconstruction

- ▶ isotropic fluorescence emission
- ▶ forward beamed direct Cherenkov light
- ▶ Rayleigh- and Mie-scattered Cherenkov light



geometry,
atmospheric corrections,
fluorescence yield



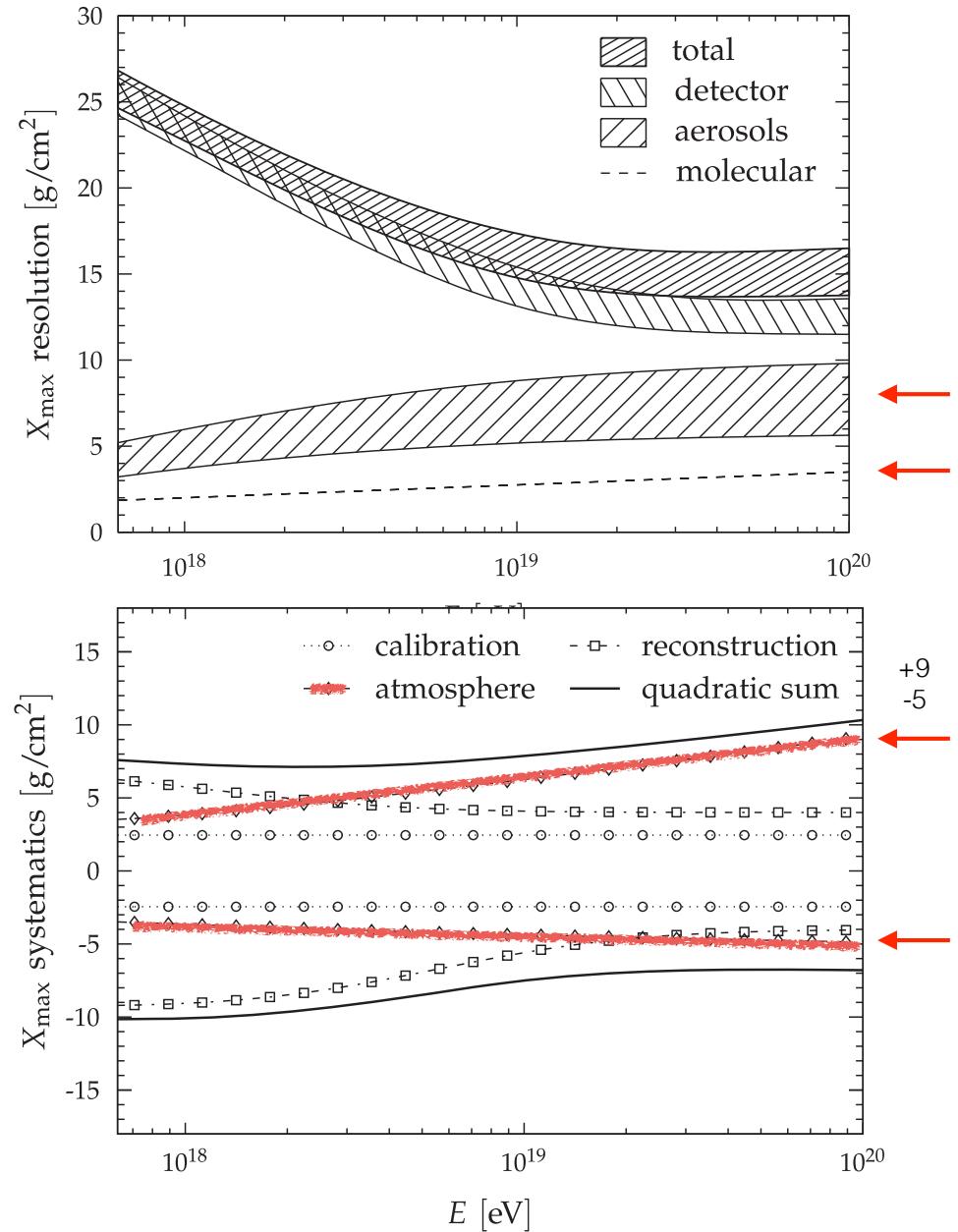
→ energy, X_{\max}

Key FD Systematic Uncertainties

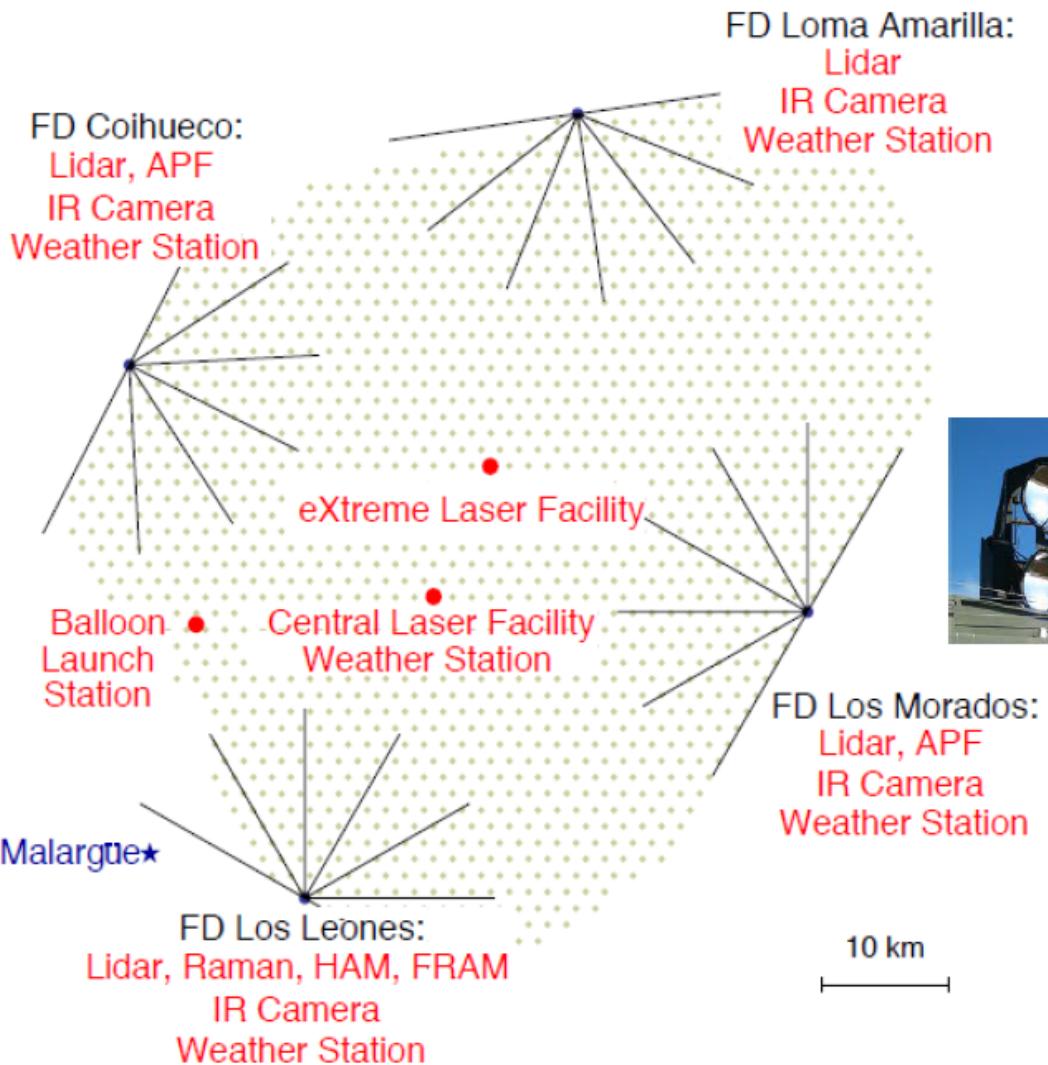
Energy (ICRC 2013)

Absolute fluorescence yield	3.4%
Fluores. spectrum and quenching param.	1.1%
Sub total (Fluorescence Yield)	3.6%
Aerosol optical depth	3% ÷ 6%
Aerosol phase function	1%
Wavelength dependence of aerosol scattering	0.5%
Atmospheric density profile	1%
Sub total (Atmosphere)	3.4% ÷ 6.2%
Absolute FD calibration	9%
Nightly relative calibration	2%
Optical efficiency	3.5%
Sub total (FD calibration)	9.9%
Folding with point spread function	5%
Multiple scattering model	1%
Simulation bias	2%
Constraints in the Gaisser-Hillas fit	3.5% ÷ 1%
Sub total (FD profile rec.)	6.5% ÷ 5.6%
Invisible energy	3% ÷ 1.5%
Statistical error of the SD calib. fit	0.7% ÷ 1.8%
Stability of the energy scale	5%
TOTAL	14%

Depth of Shower Maximum, X_{\max}

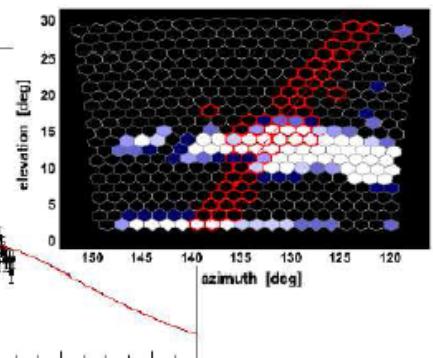


Atmospheric Monitoring Instruments

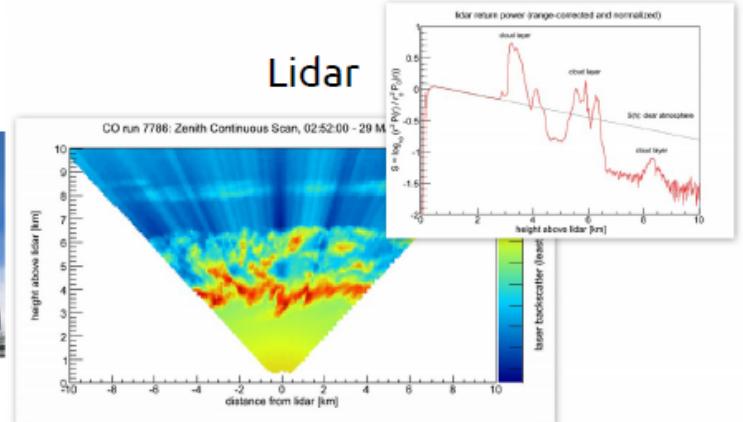


+ Global Data Assimilation System (GDAS)

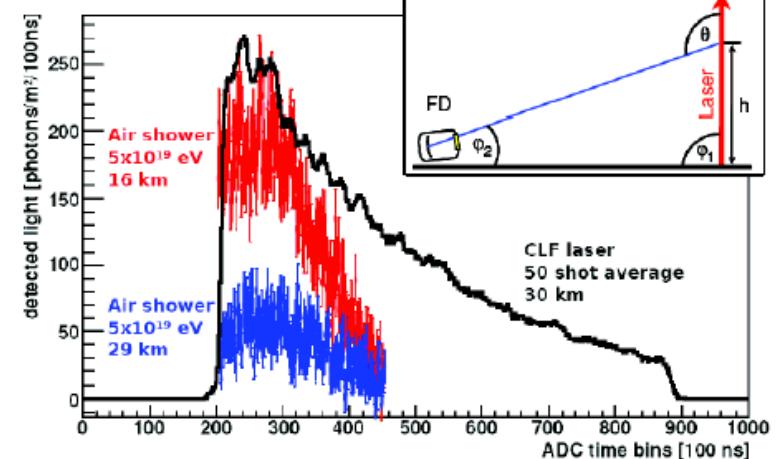
IR Cloud cameras



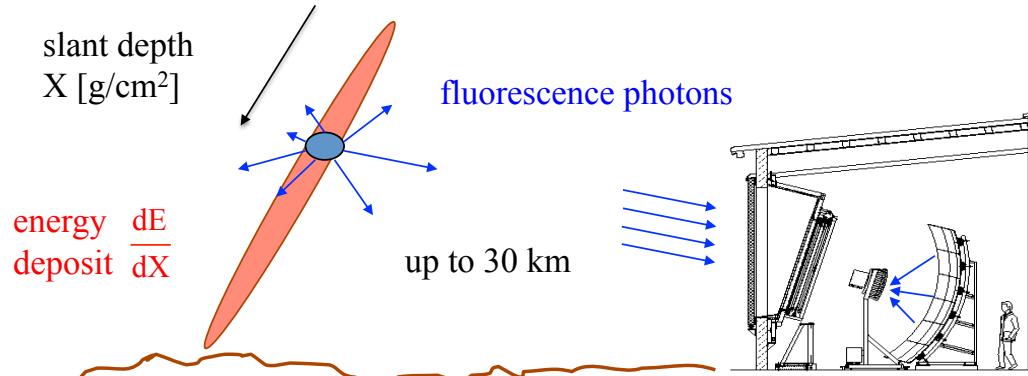
Lidar



CLF and XLF

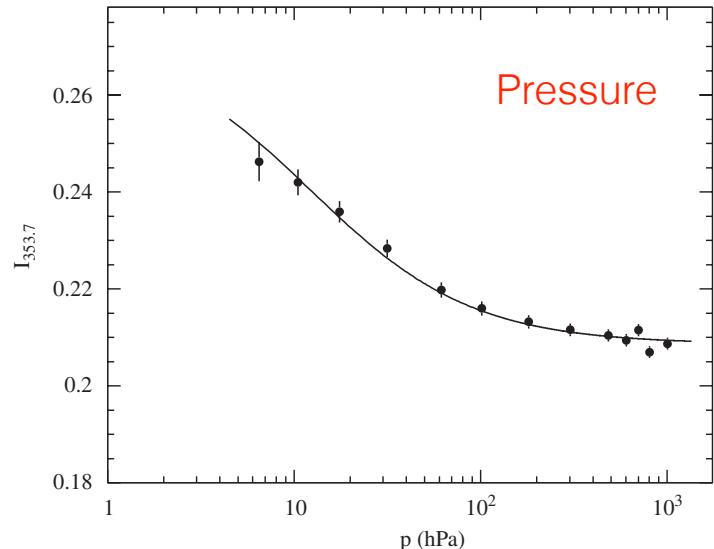


Fluorescence yield and the atmosphere

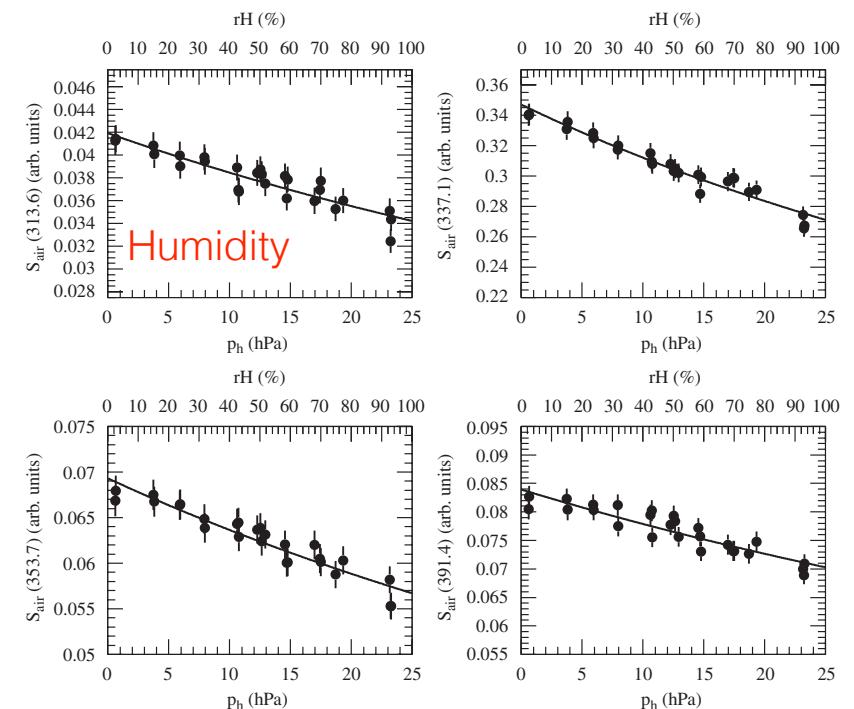
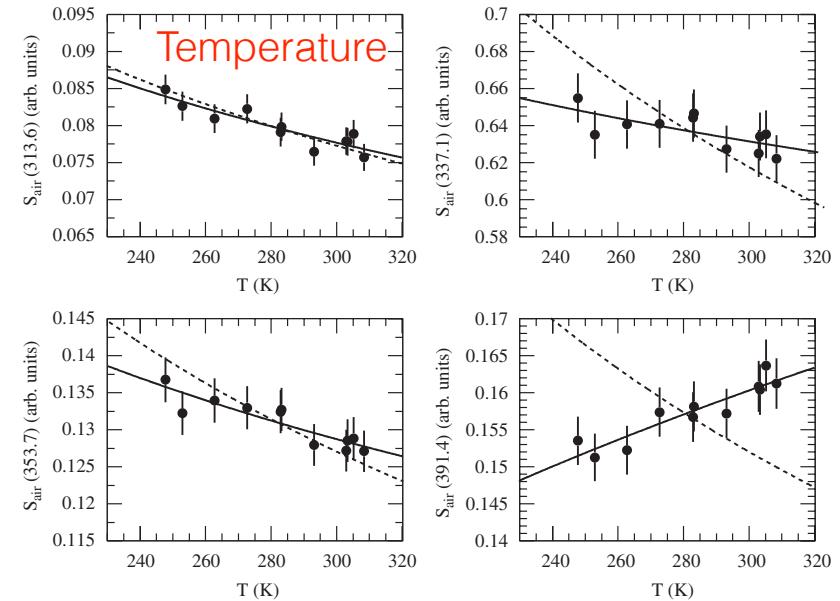


- competition between excitation and quenching
- quenching by nitrogen, oxygen and water
- temperature dependence includes collision probability & cross-sections
- AIRFLY measurements of coefficients, now much better known

Get P,T, humidity profiles from GDAS

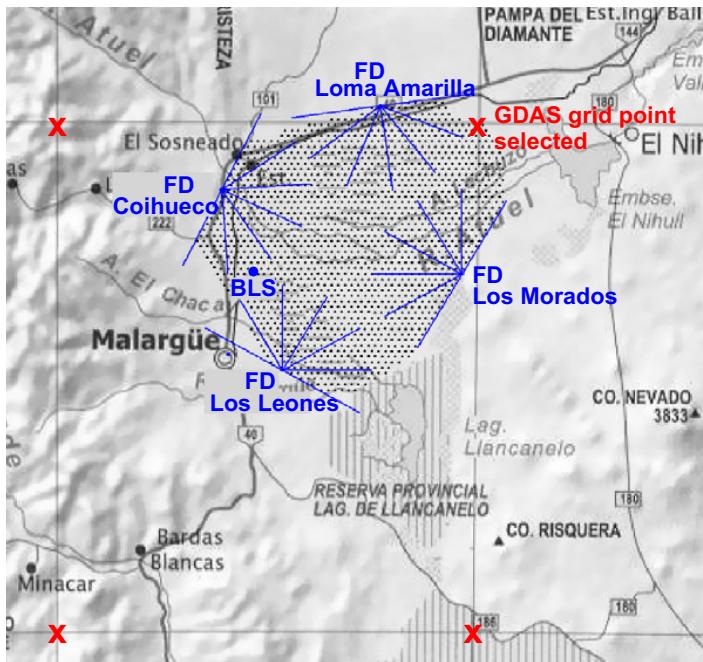


AIRFLY collaboration
NIM A597 41-54 (2008)
Astropart Phys, 42 90 (2013)

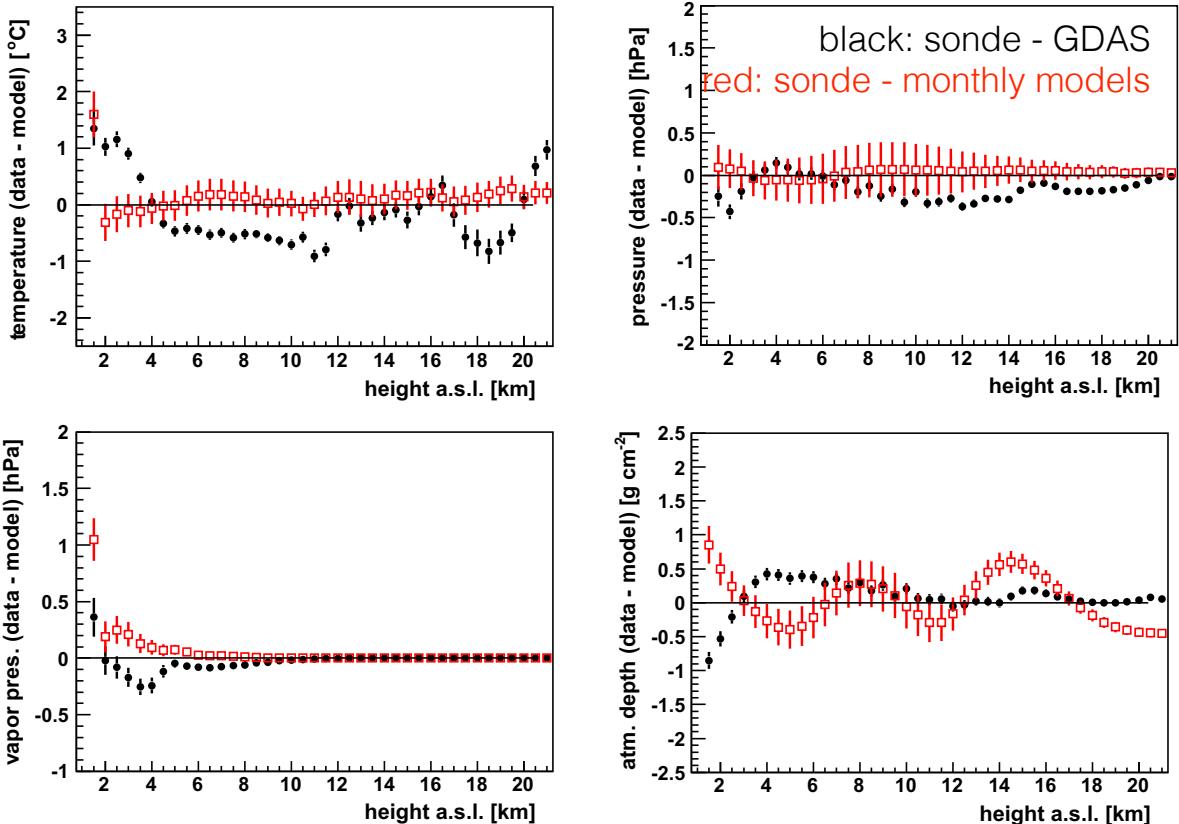


GDAS validation

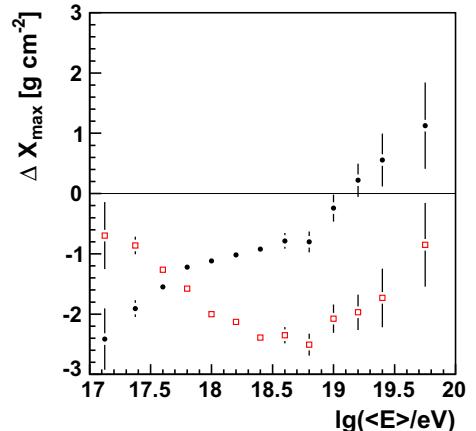
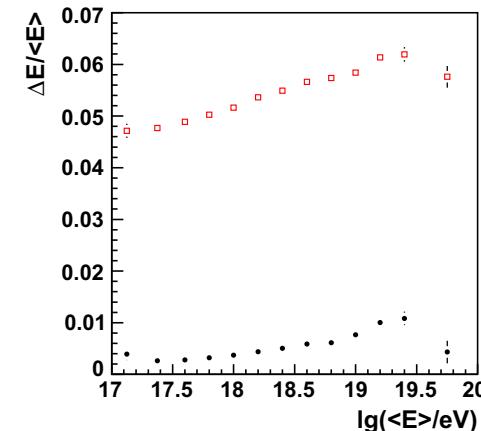
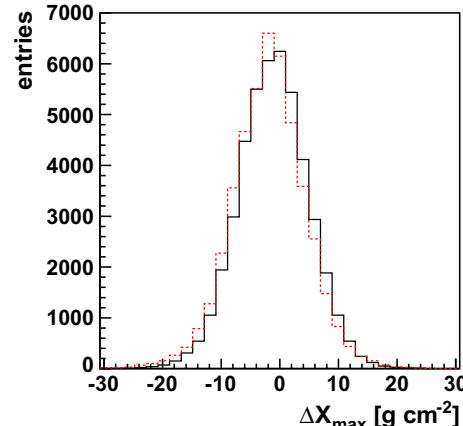
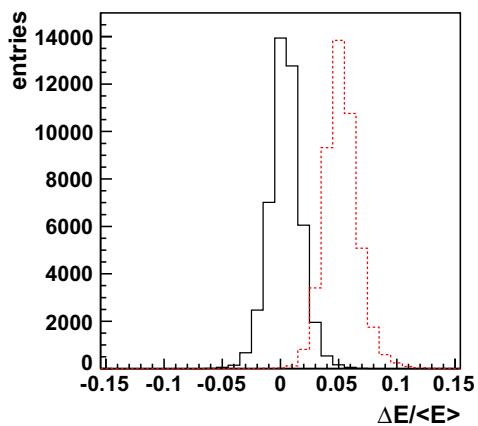
P. Abreu et al./Astroparticle Physics 35 (2012) 591–607



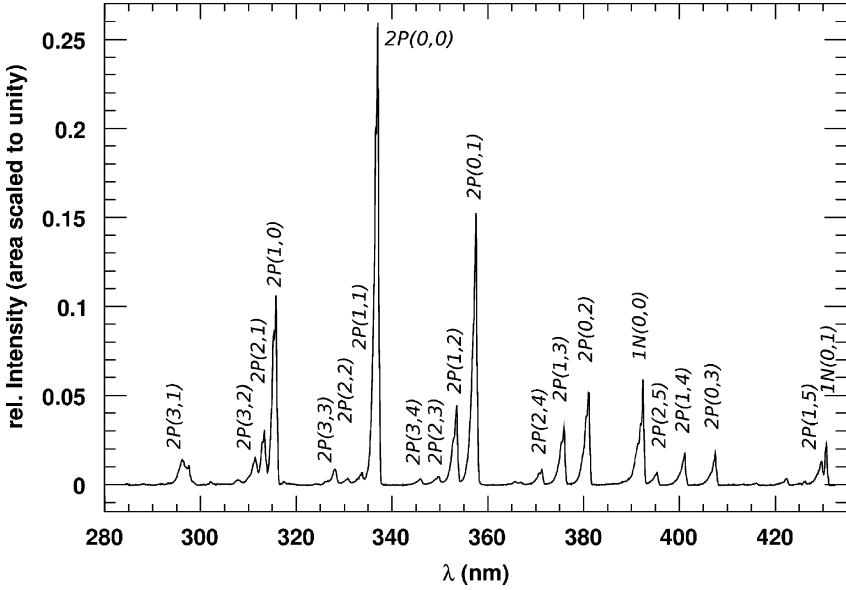
Comparing GDAS with radiosondes 2005-2008



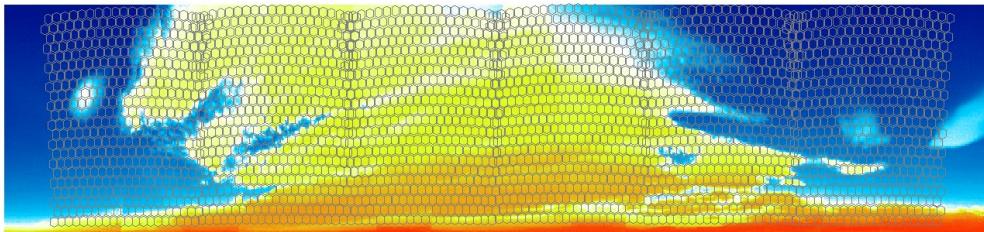
Black: effect of using GDAS rather than monthly models (ignore red)



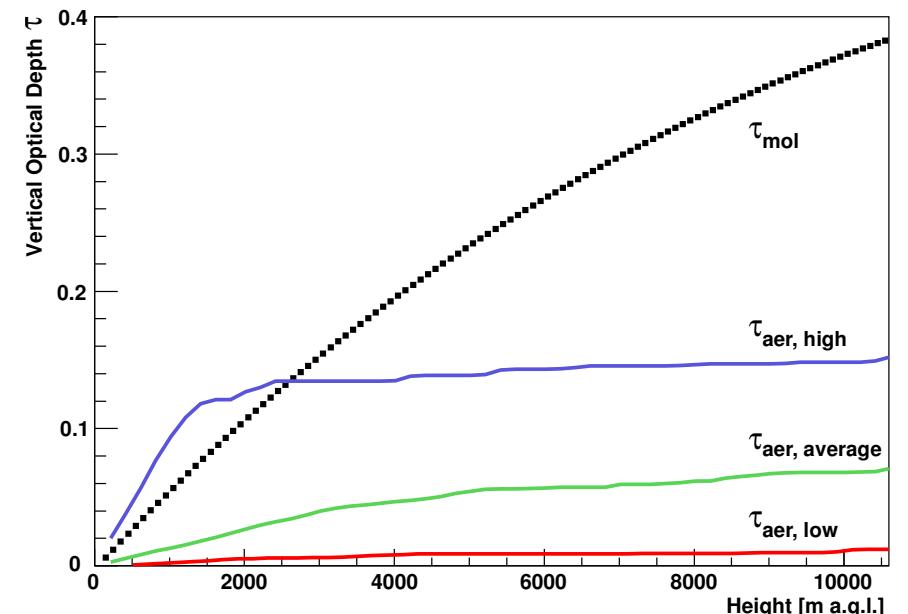
Optical attenuation



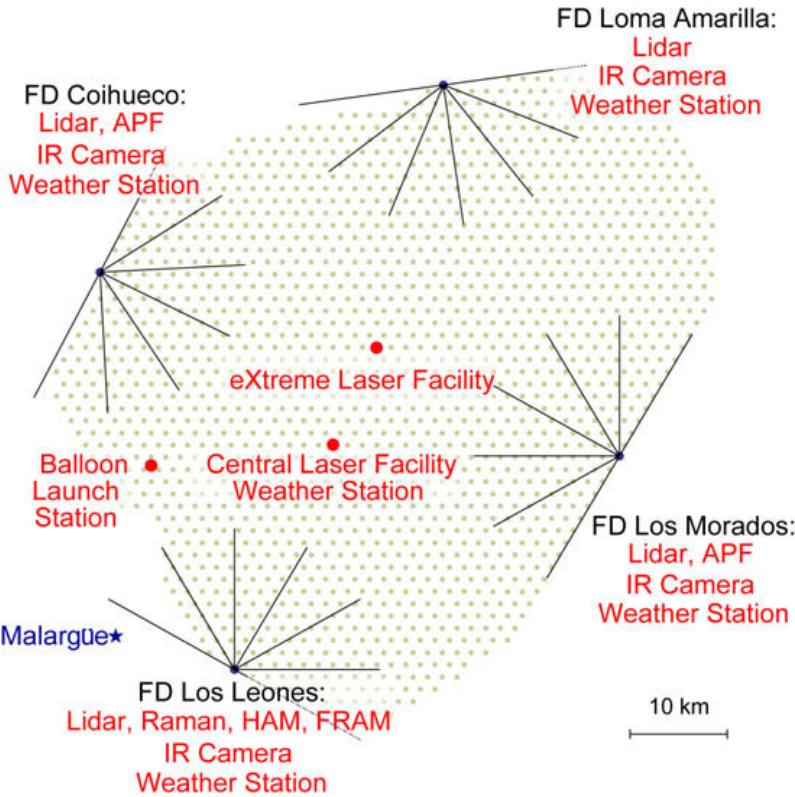
- Attenuation due to Rayleigh and aerosol scattering, and cloud
- negligible absorption (ozone)
- Molecular: GDAS
- Aerosols: Central laser facilities, Raman lidar, FRAM, (elastic lidars)
- Cloud: IR cameras, GOES satellite (10km^2), scanning elastic lidars, FRAM



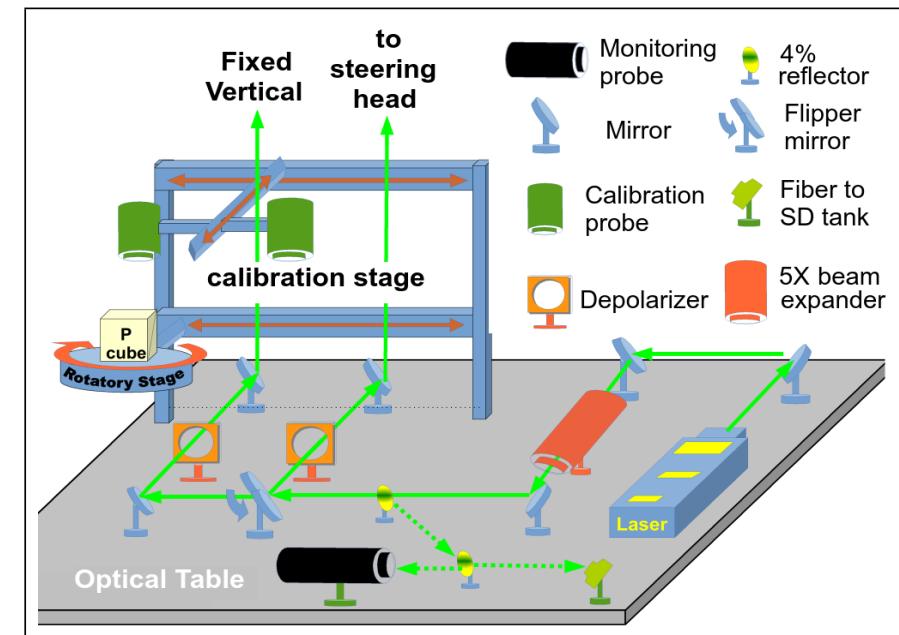
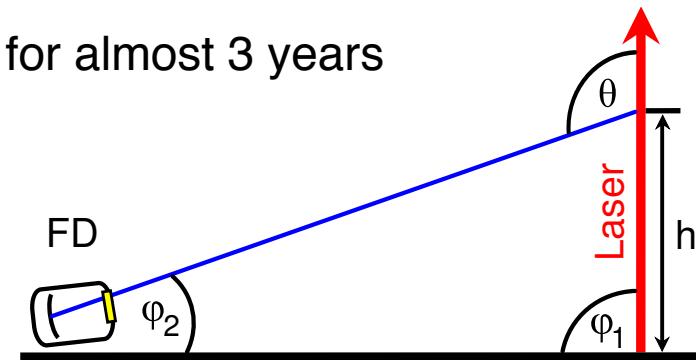
Typical molecular and aerosol VODs



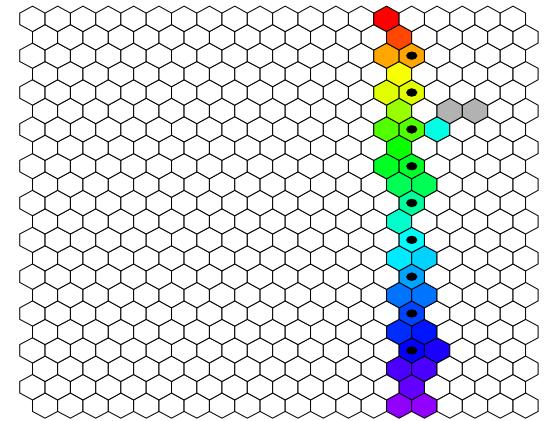
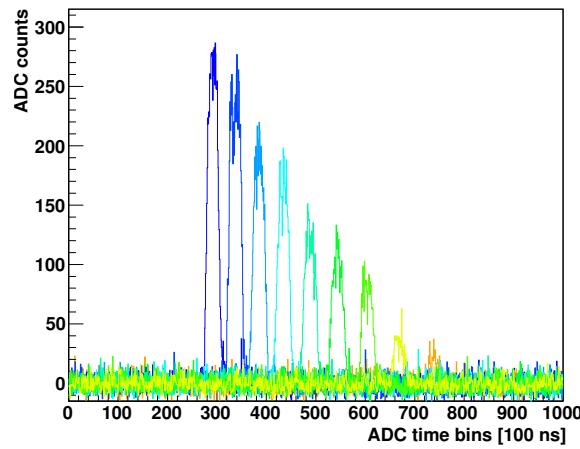
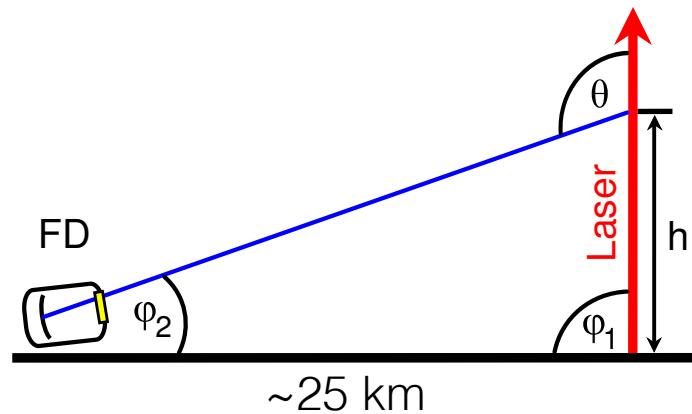
Aerosols



- CLF and XLF - *bi-static* lidar
 - two methods
 - based on reference night (assumed aerosol-free)
- (mono-static lidars used in the past)
- Raman LIDAR for almost 3 years



CLF methods for Vertical Aerosol Optical Depth τ_{aer}



- 355nm frequency-tripled YAG laser
- $\sim 6 \mu\text{J}$ per pulse
- 50 shots every 15 minutes

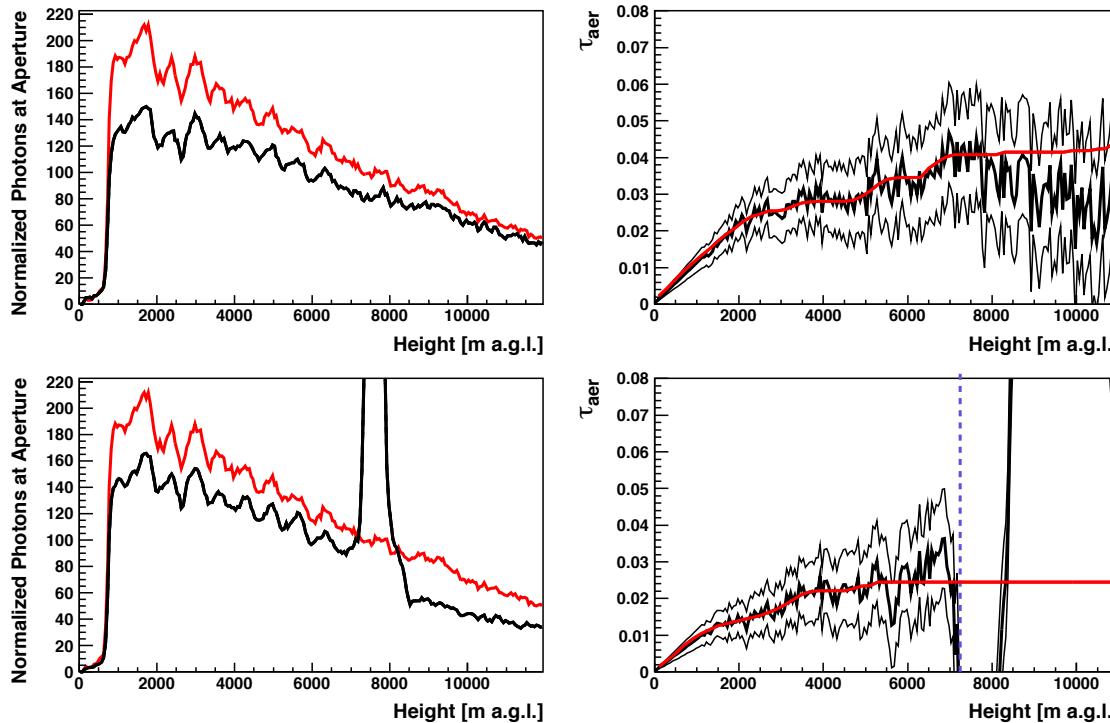
CLF methods for Vertical Aerosol Optical Depth τ_{aer}

“Data normalised” (DN) method

$$\tau_{\text{aer}}(h) = - \frac{\sin \varphi_1 \sin \varphi_2}{\sin \varphi_1 + \sin \varphi_2} \left(\ln \left(\frac{N_{\text{obs}}(h)}{N_{\text{mol}}(h)} \right) - \ln \left(1 + \frac{S_{\text{aer}}(\theta, h)}{S_{\text{mol}}(\theta, h)} \right) \right)$$

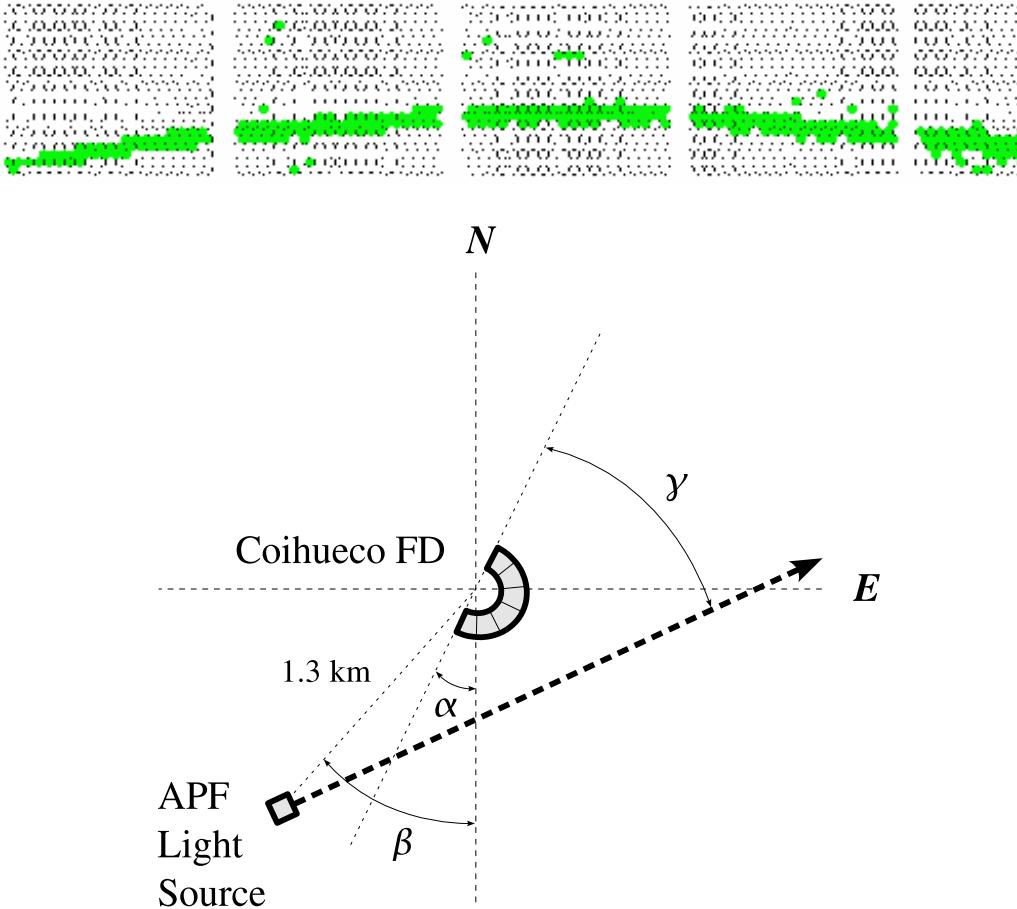
normalised to
1 μJ energy

red: reference night

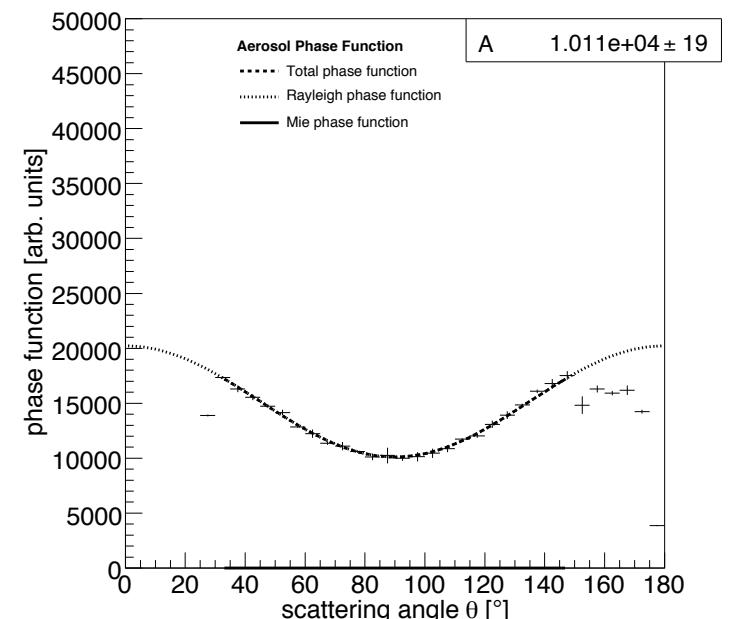
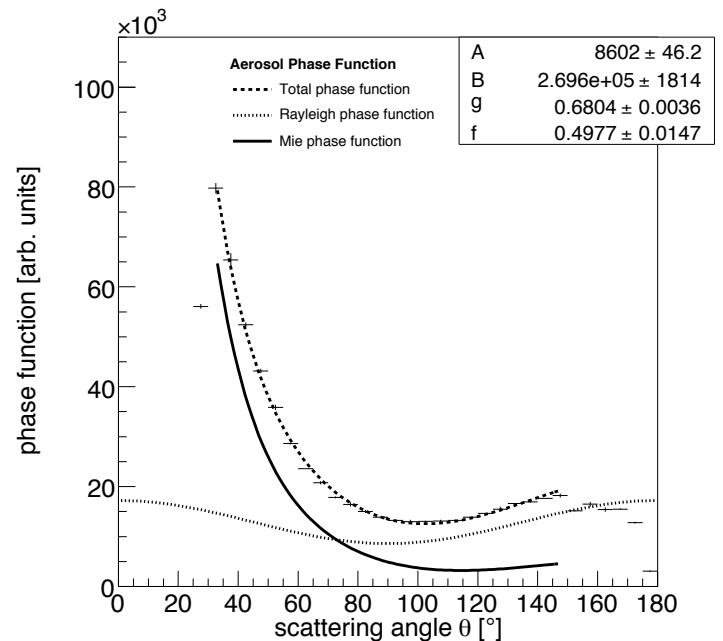


Aerosol Phase Function (APF) Monitor

(one use - checking reference nights)

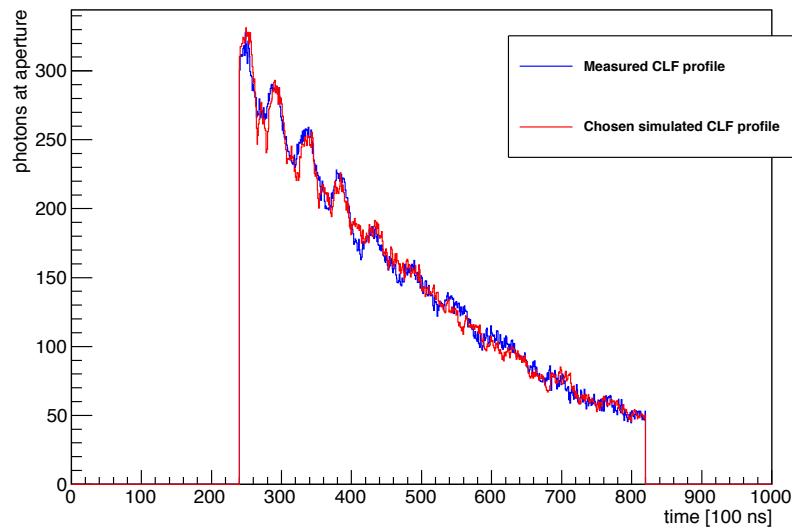
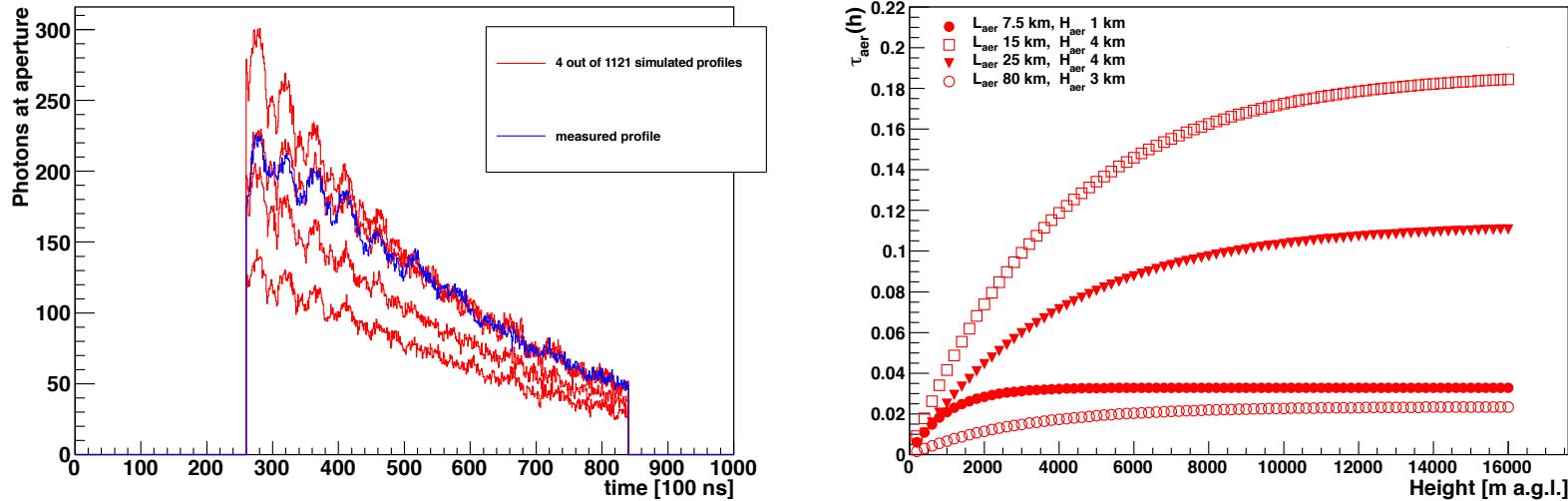


Near-horizontal collimated xenon flash lamps
at two FD sites



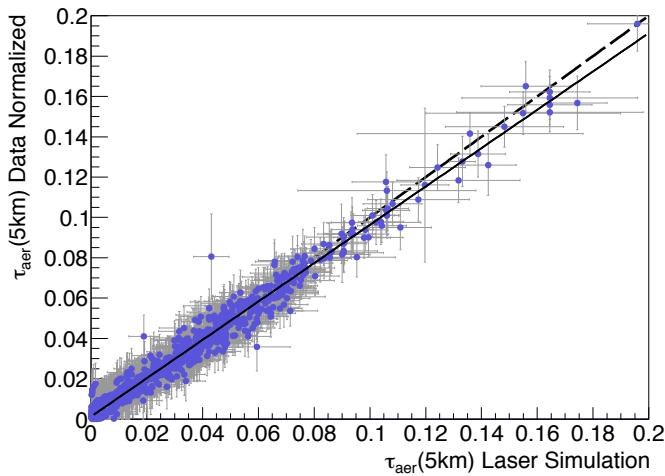
CLF methods for Vertical Aerosol Optical Depth τ_{aer}

“Laser Simulation” (LS) method

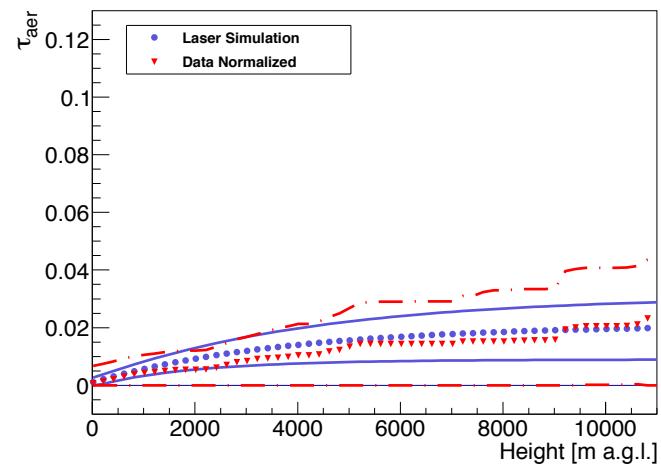


For technical reasons (normalisation of simulations),
this method also requires use of a reference (clear) night.

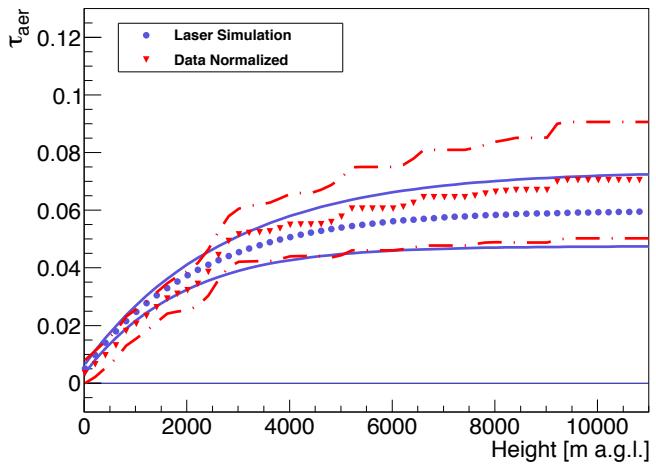
CLF methods for Vertical Aerosol Optical Depth τ_{aer}



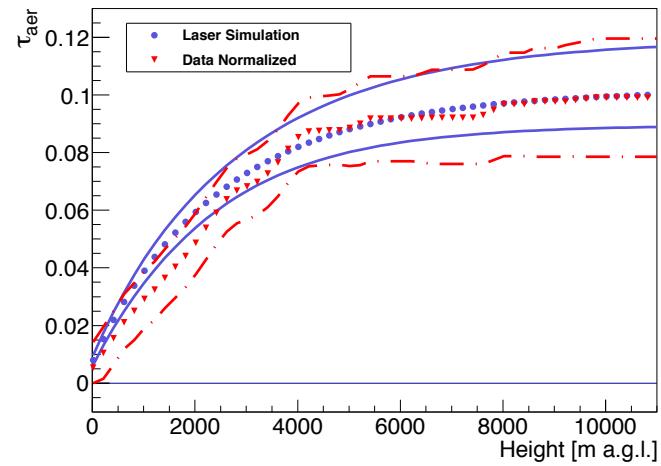
(a) Correlation between the analyses.



(b) Low aerosol attenuation.

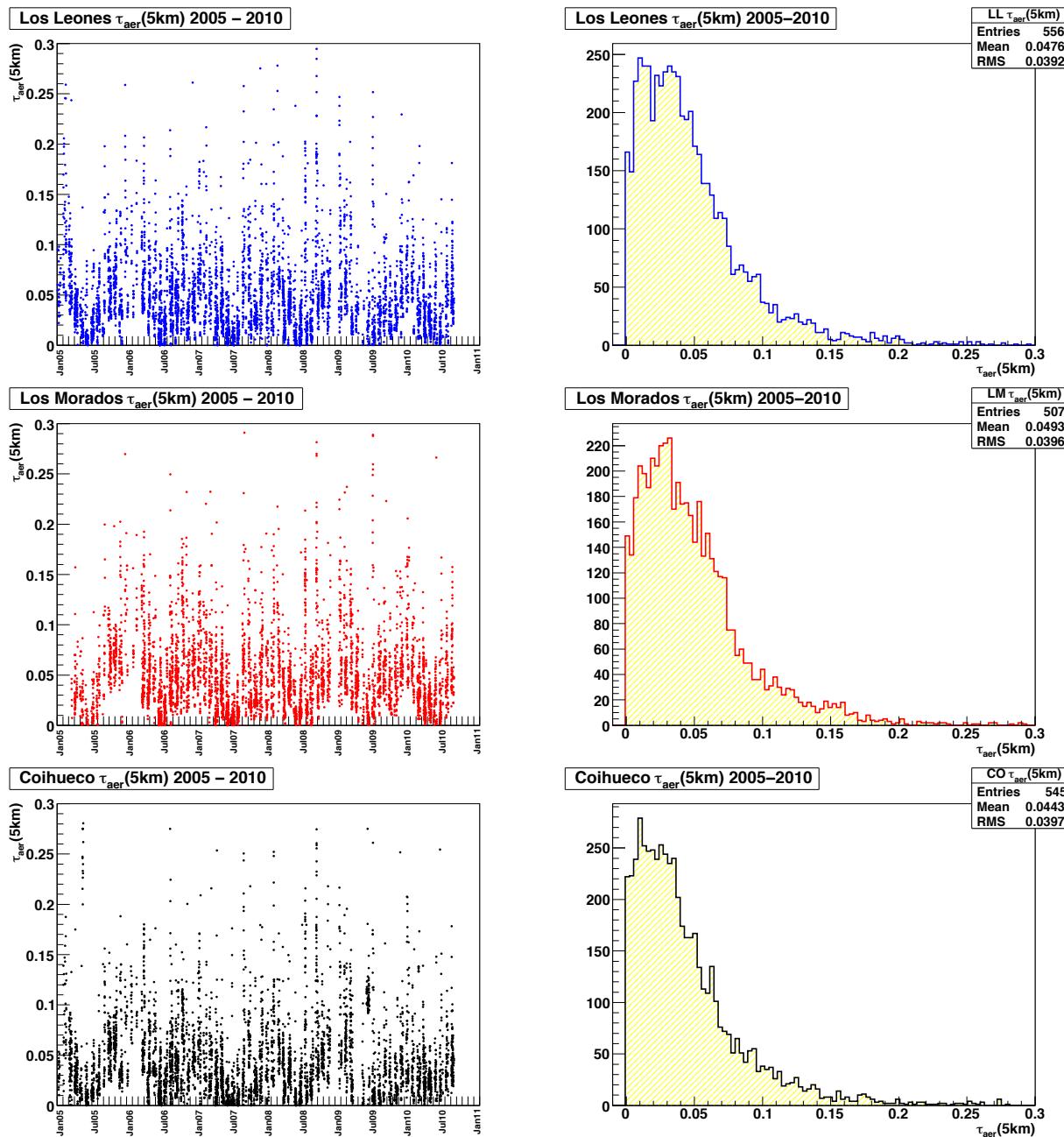


(c) Average aerosol attenuation.

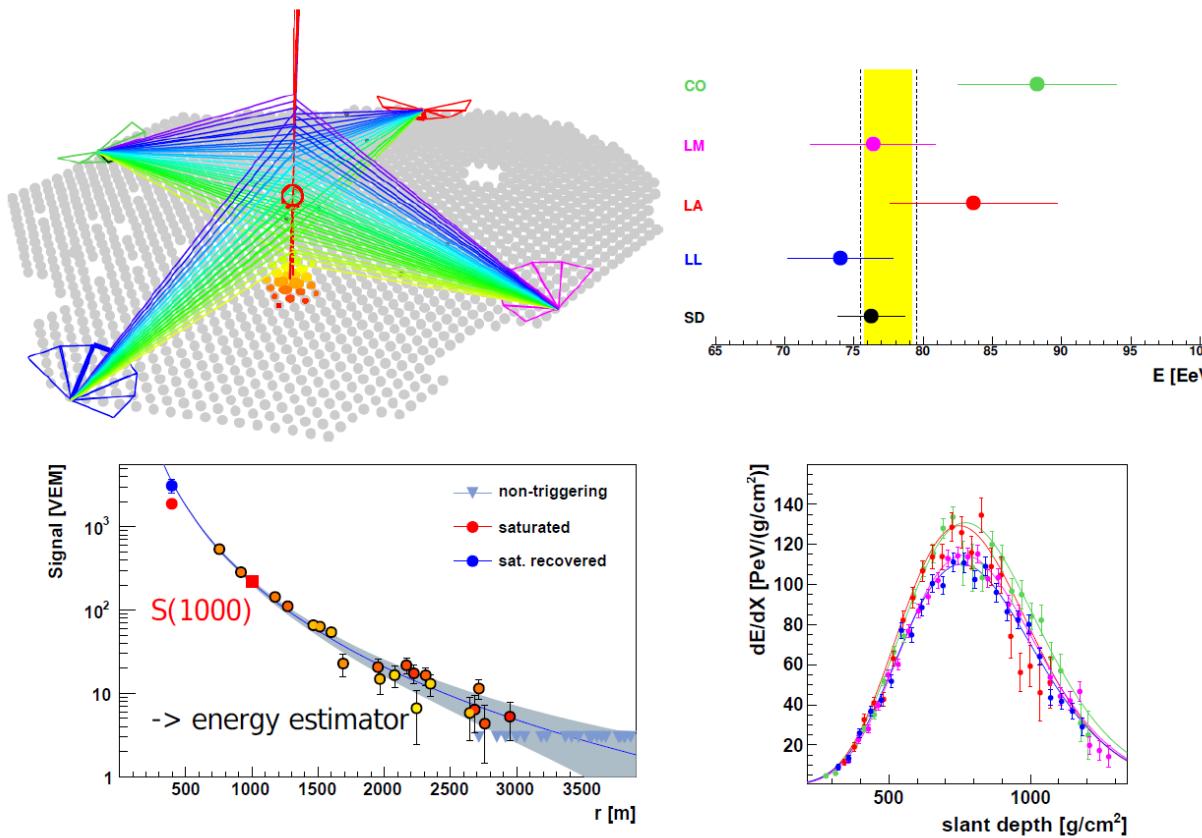


(d) High aerosol attenuation.

CLF methods for Vertical Aerosol Optical Depth τ_{aer}



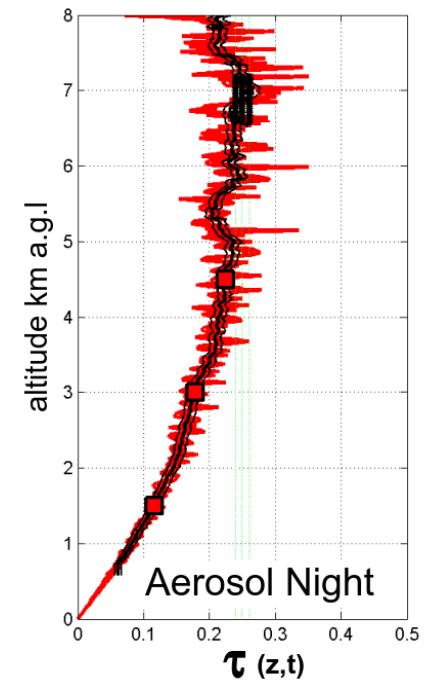
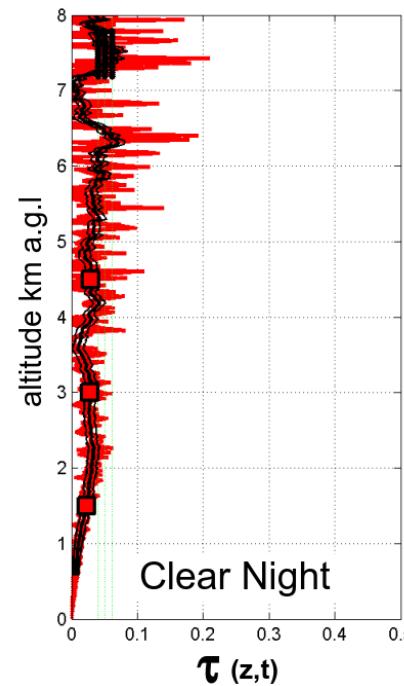
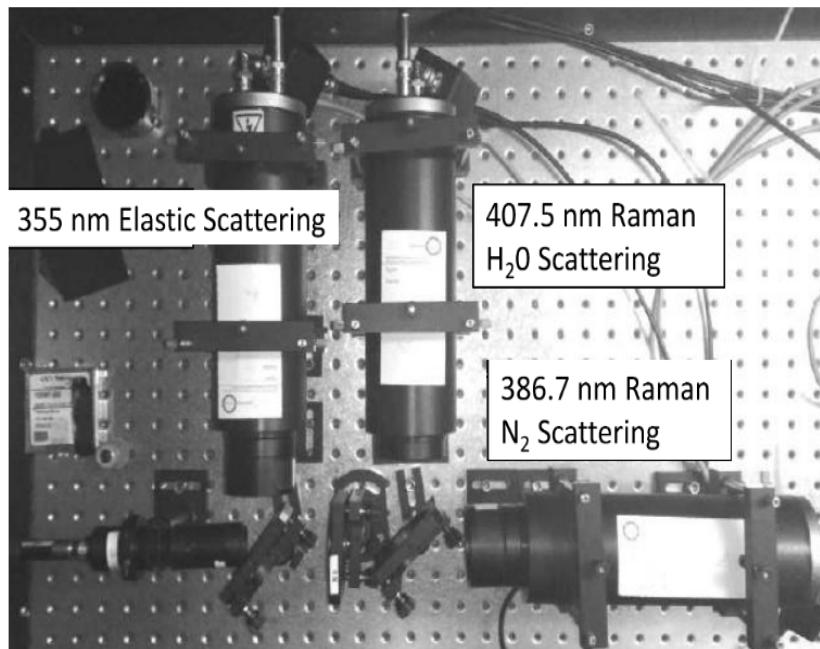
Cross-checks



- We exploit stereo-views of showers to check assumptions of atmospheric attenuation (HiRes “stereo-balance”)
- Even hybrid events can be used to check E_{FD}/E_{SD} as a function of the distance to the FD.

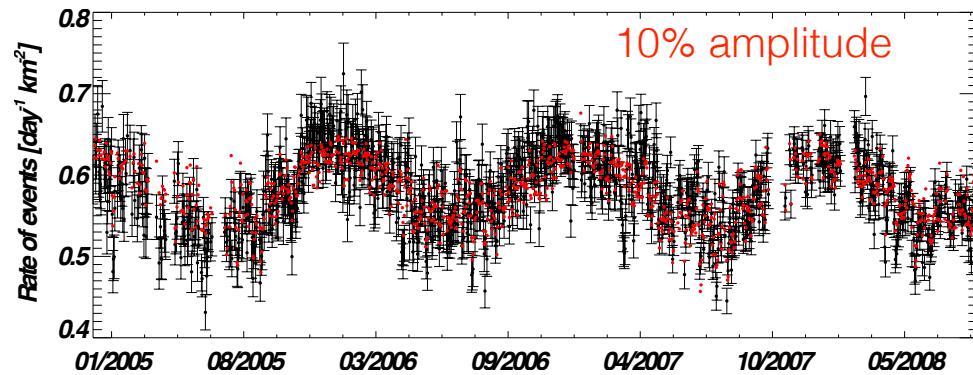
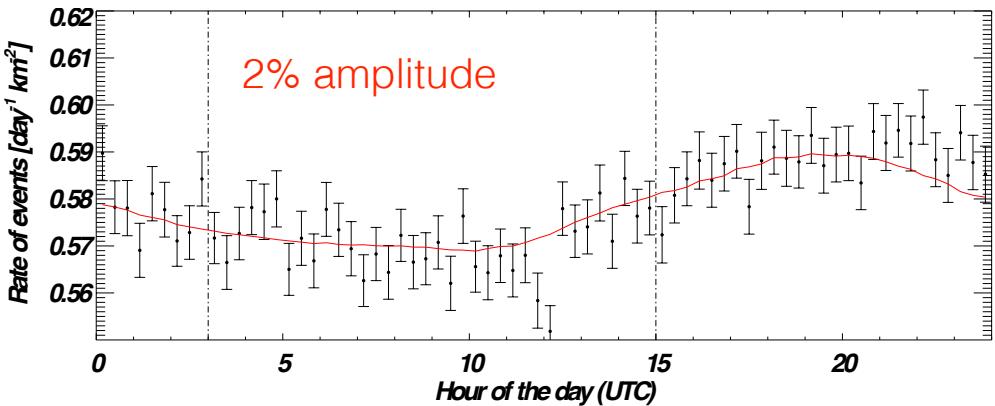
Cross-checks

Raman LIDAR receiver
(since Nov. 2013 at CLF site)



- uses 355nm YAG at CLF
- runs 3 times per night (12 minute integration time)
- running reliably, comparisons in progress

Surface detector energy - P,T dependence



- Shower rates above a given assumed energy threshold show diurnal and seasonal variations
- due to
 - density ($\rho \sim P/T$) related changes to the Moliere radius (most important)
 - shower attenuation with increasing P
- can be corrected empirically using weather station data (5 min latency)
- particularly important in broad scale anisotropy studies, where harmonic amplitudes are $\sim 1\%$

Conclusions

- the challenge of monitoring the atmosphere over 3000 square kilometres is achieved with a suite of instruments
- measurement redundancy is important!
- creative use of air shower data for cross-checks is fruitful
- work continues to improve understanding, cross-check our assumptions and results

Photo: Steven Saffi, University of Adelaide

